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# Using AIRS to study dust and cirrus

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- AIRS launched in May 2002 on board NASA-Aqua (A-train); operational since Sept 2002
- AIRS is a hyperspectral infrared sounder
- AIRS has low noise, high resolution thermal IR channels (650 - 2800  $\text{cm}^{-1}$  with  $\nu/\delta\nu \simeq 1200$ ) (3.7 to 15.4  $\mu\text{m}$ )
- 13.5 km footprint, scans  $\pm 45\text{deg}$  from nadir, twice daily global coverage
- Produces temperature profiles with 1K/km accuracy (rms errors  $\leq 1.25\text{K}$ ), water vapor (rms errors 20-40%) and trace gas profiles
- Very stable instrument with well characterized radiances
- Standard products include cloud top heights

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- Magnitude of climate forcing by clouds/aerosols is uncertain, especially in the longwave
- Can use MODIS to identify dusty scenes, MISR to obtain optical depths and CERES to obtain broadband TOA LW flux
- AIRS has sensitivity to silicate-based absorbers and cirrus : radiance spectra exhibit their spectral signatures
- Can potentially use AIRS to study all three over ocean or land eg use AIRS radiances day and night, over ocean and land to
  - detect dust
  - retrieve optical depths
  - obtain quick estimates of OLR forcing

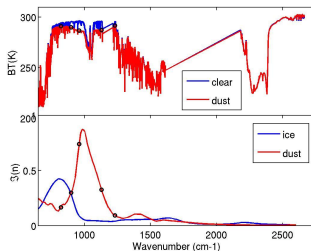
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- Set up a sequence of “threshold dust cloud tests”
- 5 channels chosen are [822.4 900.3 961.1 1129.03 1231.3]  $cm^{-1}$
- Tests involve split window brightness temperature differences
- Use  $t=380$  over water;  $t=360$  over land (warning : needs improvement)

# Long Range Transport of Sahara Dust

## AIRS data for July 2003

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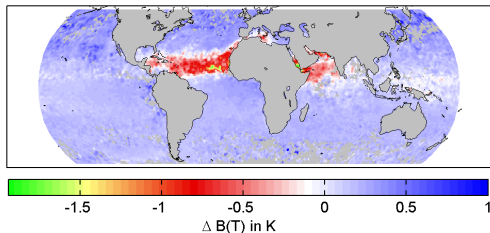
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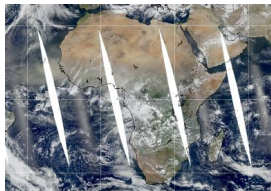
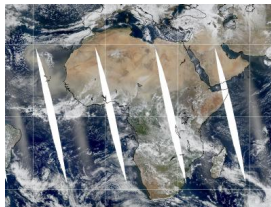
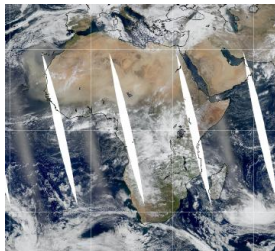
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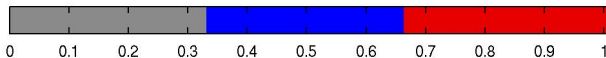
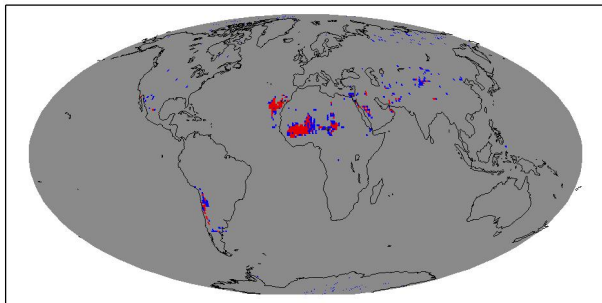
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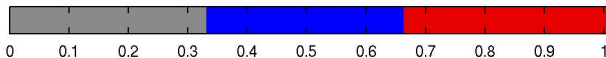
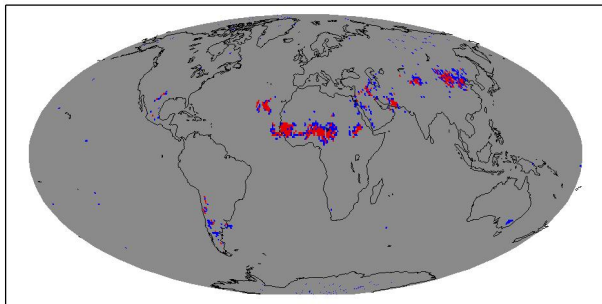
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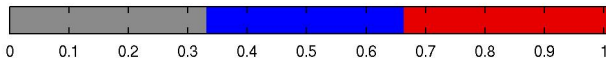
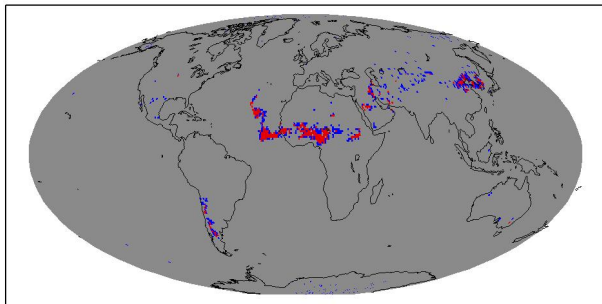
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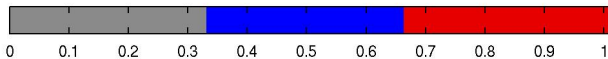
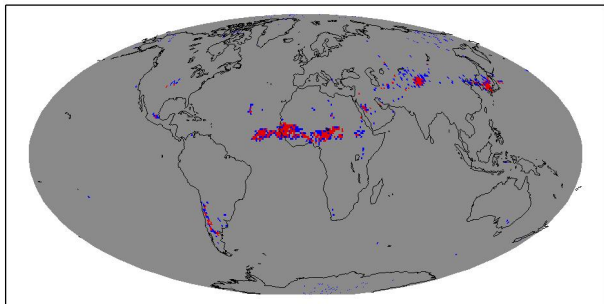
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- use SARTA (PCLSAM : Chou et al, AMS Jan 1999 pg 159)
- uses Masuda emissivity for ocean
- uses Global Infrared Land Surface Emissivity Database (SSEC/U.Wisc) (E. Borbas, S. Wetzels-Seemann, R. O. Knuteson, P. Antonelli, J. Li and H.-L. Huang)
- uses ECMWF (or AIRS retrievals) for  $T(z), Q(z)$  fields, with adjusted surface temperature (George Aumann) for sea and land
- very fast  $\leq 1$  second per profile (even if looping over  $ptop, dme$ )

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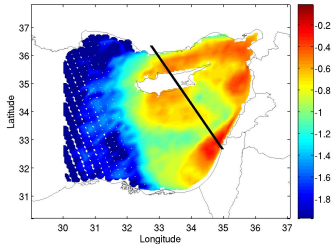
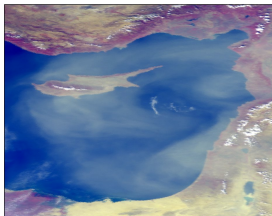
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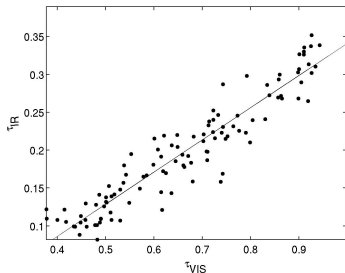
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MODIS channel 2 (0.55  $\mu\text{m}$ ) compared to AIRS 900  $\text{cm}^{-1}$

$\tau_{IR} = 0.425\tau_{VIS} - 0.084$ , with a correlation of 0.935 GRL Paper with Sergio

**DeSouzaMachado, Larrabee Strow, Scott Hannon and Howard Motteler**

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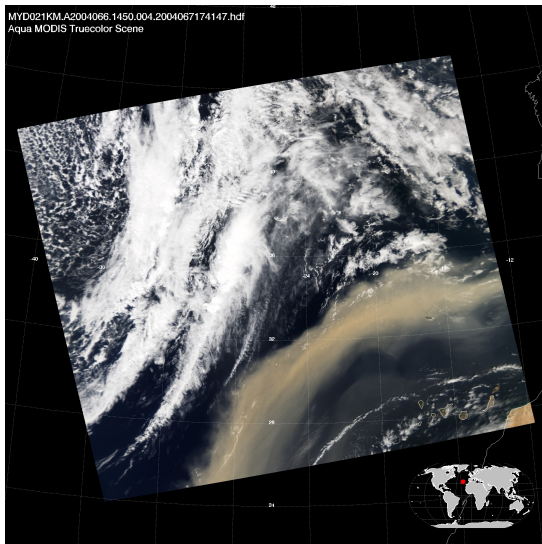
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# True color image made from MODIS data, for March 6, 2004 at approximately 1430 UTC



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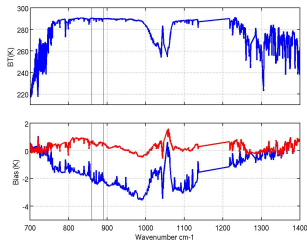
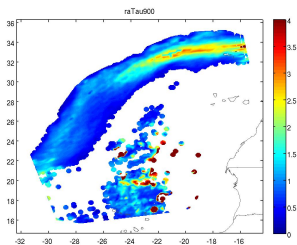
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Optimum particle diameter  $\simeq 2 \mu\text{m}$ ; optimum height  $\simeq 1 \text{ km}$

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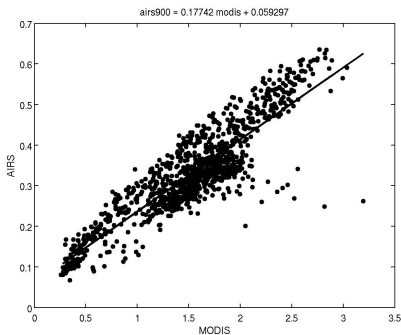
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AIRS infrared optical depths at  $900 \text{ cm}^{-1}$  plotted against MODIS Ch 2 (550 nm) visible optical depths, for dusttop at 600 mb. At 900 mb (1.0 km),  $\frac{\tau_{\text{AIRS}}}{\tau_{\text{MODIS}}} \approx 0.5$



# MODIS image of duststorm on March 3, 2004 over N.W.Africa

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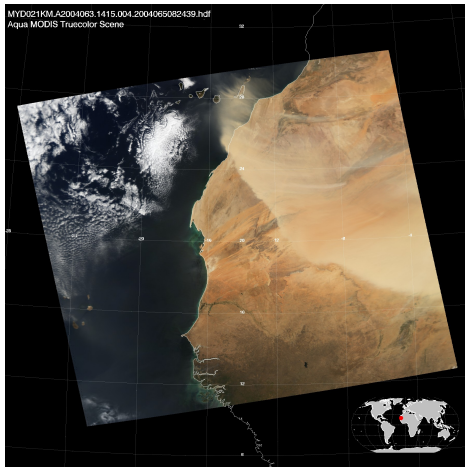
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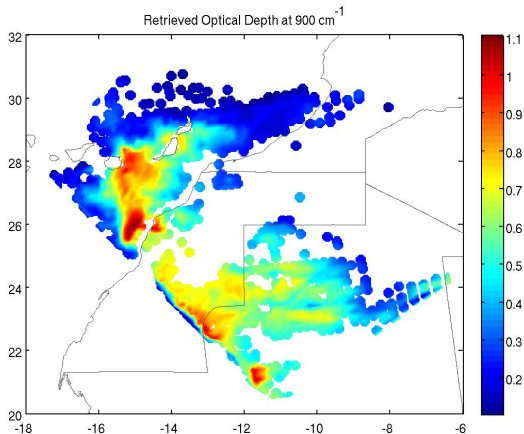
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# Infrared Retrievals from many global duststorms (over ocean)

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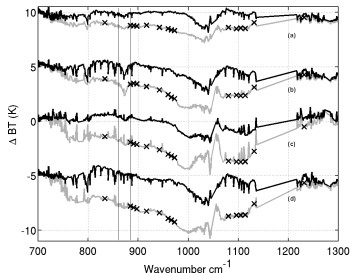
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- (a) Libyan/Egyptian coast (02/28/2005)
- (b) Eastern Mediteranean (10/19/2005)
- (c) China Sea (11/12/2002)
- (d) W. African coast (07/25/2004)
  - All show the “V” shape in  $800\text{--}1200\text{ cm}^{-1}$  (silicate absorber)
  - Notch feature between  $860\text{ and }880\text{ cm}^{-1}$  is strongest in *b, c*

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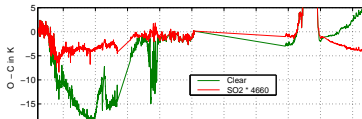
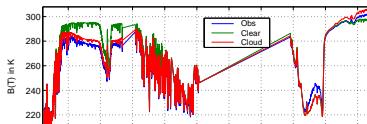
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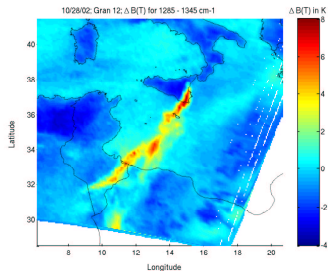
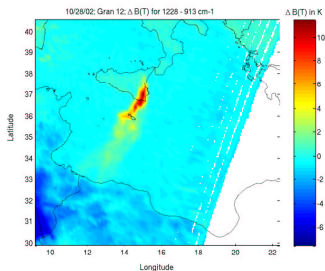
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Oct 28, 2002; Granule 123; Profile 1502





**GRL Paper with Simon Carn, Yvonne Edmonds, Larrabee Strow, Sergio DeSouzaMachado, Scott Hannon and Howard Motteler**

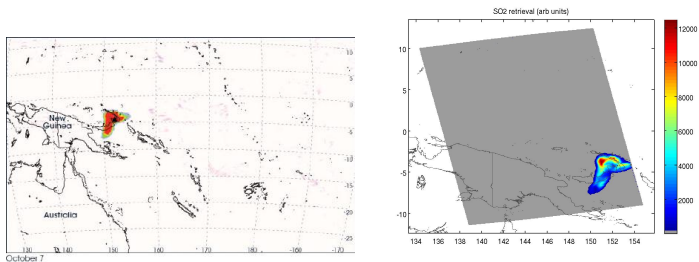
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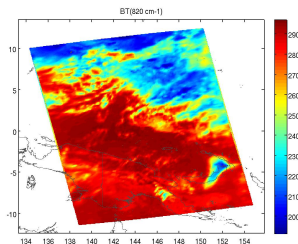
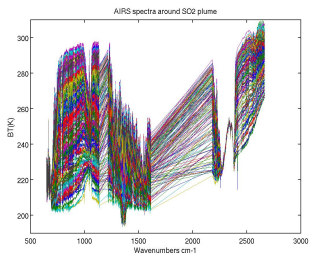
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LHS : <http://earthobservatory.nasa.gov/Newsroom/NewImages/>  
RHS : Thanks to Scott Hannon! SO<sub>2</sub> multiplier ( $\times$  standard SO<sub>2</sub> profile)



Thanks to Scott Hannon!

LHS shows AIRS spectra in the SO<sub>2</sub> plume region; notice strong SO<sub>2</sub> signal and cirrus contamination

RHS shows BT(820 cm<sup>-1</sup>) and shows cirrus cloud in vicinity of SO<sub>2</sub> plume

Using the PCLSAM model (Chou et al, AMS Jan 1999 pg 159) can reparameterize optical depth  $\tau$  with atm gases only to  $\tau \rightarrow \tau(atm) + \tau(scatter, E, \omega, g)$

Radiance at the top of a cloudy sky atmosphere

$$R(\nu) = \epsilon_s B(\nu, T_s) \tau_{1 \rightarrow N}(\nu, \theta) + \sum_{i=1}^{i=N} B(\nu, T_i) (\tau_{i+1 \rightarrow N}(\nu, \theta) - \tau_{i \rightarrow N}(\nu, \theta))$$

This is same as clear sky OLR equation, and so can compute estimates of OLR forcing



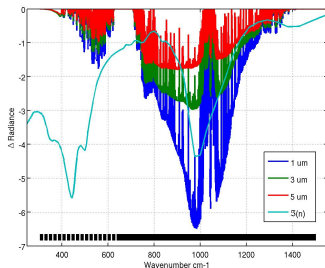
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Simulated thermal IR and Far IR changes in TOA radiance, for realistic dust cloud of particle diameters 1,3,5  $\mu\text{m}$ , with the refractive index (scaled) shown in cyan. 90% of the OLR forcing comes from the region covered by the AIRS thermal IR channels (shown in solid black). Vertical units are in  $\text{mWm}^{-2}\text{sr}^{-1}/\text{cm}^{-1}$ .

# MODIS image of duststorm on March 3, 2004 over N.W.Africa

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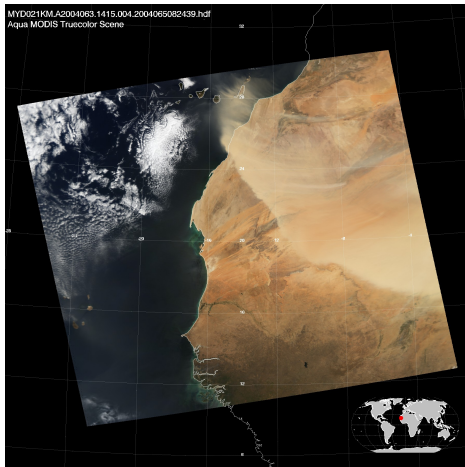
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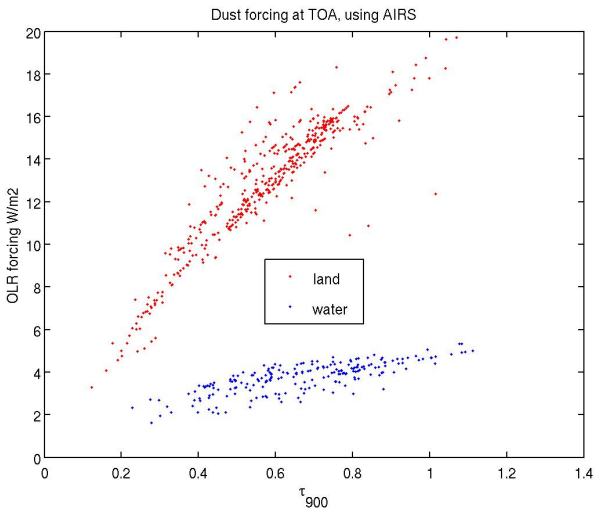
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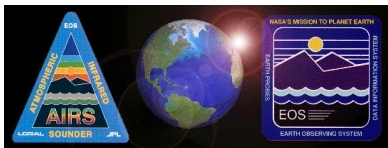
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Ability of hyperspectral IR sounders to detect dust, retrieve optical depths and compute lower bound estimates of OLR forcing



<http://earthobservatory.nasa.gov/>  
<http://www-airs.jpl.nasa.gov/>  
<http://asl.umbc.edu/>  
[sergio@umbc.edu](mailto:sergio@umbc.edu)

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